Analysis of Streamflow Depletion and Well Interference under Various Conditions

In the process of revising the County well ordinance and developing policies to minimize impact on streamflow and public trust values, County staff have analyzed the potential effects of individual wells under various conditions using several different analytical models. These models are used to provide a sensitivity analysis and evaluate the extent to which different factors may influence streamflow depletion in Santa Cruz County. However, it's important to note that analytical models rely on various assumptions, commonly including the presumption of steady-state conditions for the stream and aquifer. In reality, the degree of stream depletion is likely to fluctuate in response to changing climate conditions over time. Modeled estimates of depletion are likely somewhat inaccurate as the environment of the Santa Cruz Mountains is inconsistent with many of the underlying assumptions upon which the models are based.

The variability of Santa Cruz County's climate, geology, topography, ecological, and stream conditions makes the establishment the thresholds for each tier, and allowed additional stream depletion for specific streams is challenging. The thresholds and limits are formed on empirical data, current and expected water use conditions, model simulations, and expert judgment, and are subject to refinement as new data and models emerge. It's important to recognize that model-derived values are especially influenced by specific hydrogeologic conditions unique to our community. Therefore, while our approach provides a valuable framework for assessing local impacts, its application to other regions should consider local conditions and professional discretion.

In addition, the amount of total depletion that is estimated to be presently occurring based on numeric groundwater models and flow measurements (Table 4) is considerably less than the amount that would be calculated by multiplying the number of current wells by the worst-case calculations of the direct effect of individual wells provided by the analytical models.

Estimates of streamflow depletion were calculated and analyzed using a combination of models including the USGS web based calculation, STRMDEPL08 (available at https://mi.water.usgs.gov/software/groundwater/CalculateWell/index.html), the analytical depletion function (ADF) model developed by Li et al. 2022 (found at: https://github.com/FoundrySpatial/streamDepletr), and ADF model developed by Bakker in 2013 (found at: https://github.com/mbakker7/ttim). Below is a summary of the key results, along with their policy implications, detailed observations, a more in-depth discussion of our sensitivity analysis and modeling tools, and appendices providing supporting documents of our analysis.

Summary and Policy Implications:

In our sensitivity analysis, we examined various aquifer conditions and potential well mitigation strategies to account for a wide range of effects. This included assessing both unconfined and confined aquifers, incorporating detailed regional data on the variability in aquifer properties and confining layers. Specifically, we analyzed the Santa Margarita Formation (Tsm) under unconfined conditions due to its wide range of aquifer properties (e.g., hydraulic conductivity values from 2 to 130 ft/day), aiming to establish upper and lower bounds for streamflow depletion estimates. Similarly, the Monterey Formation (Tm) was studied to understand streamflow impacts under confined conditions.

In general, stream depletion impacts are more significant in interconnected wells within aquifers of high hydraulic conductivity and low storage coefficients, and less pronounced in aquifers with low conductivity and high storage coefficients. Our analysis shows that in unconfined aquifers without well seals, stream depletion correlates closely with extraction rates over relatively short distances and time periods (about 700 days). For example, pumping from the Santa Margarita Formation in an unconfined state can deplete streams by up to 98% of the pumped volume.

For confined settings, over shorter time periods (approximately 2 years), wells extracting from a confined aquifer with median Tsm hydraulic conductivity values, and a confining layer consisting of median Tm hydraulic conductivity values, the estimated stream depletion is approximately 25% of the pumped volume. However, over 10 years under the same conditions, total stream depletion can rise to about 55% of the pumping volume, indicating delayed impacts, especially pronounced in confined settings. Therefore, Tier 3 permit applications must evaluate stream depletion impacts over a 10-year timeframe. For detailed hydraulic properties used and modeled results, refer to the plots in Appendix A.

We also analyzed potential well mitigation strategies, including the use of deeper well seals. Analytical models indicate that well seals are effective in reducing stream depletion, especially over short distances and periods (~200 days) when wells are within 800 feet of streams (Appendix D). For example, in an unconfined aquifer with median Tsm properties, a well with a 100-foot seal located 100 feet from the stream can reduce depletion by 60% compared to a well without a seal over 200 days. Extending the time period to 10 years reduces the effects of diminished stream depletion, yet it still can be 20 to 70% more effective than wells without a seal (Appendix B). When evaluating the impacts of wells situated at farther distances, such as 1000 feet from the stream, the effectiveness of the seal diminishes significantly (Appendix E).

Based on this analysis, we are requiring a minimum 100-foot well seal for Tier 1 applicants within 1000 feet of a stream, and a minimum 200-foot seal for Tier 2 and 3 wells within 2000 feet of a stream.

Other well mitigation strategies and their impact on stream depletion were also analyzed, including the effects of setbacks. Analytical models suggest that increasing the distance from the well to the stream can lead to minor to substantial reductions in stream depletion. For example, at 50 feet from the stream, depletion reductions range from 2 to 10%, and at 100 feet, reductions add some additional margin of protection at a range from 3 to 20%. A 1000-foot setback can reduce depletion by approximately 15 to 75%, and at 2000 feet, up to 95%. Actual impacts in Santa Cruz County can vary widely due to abrupt changes in topography, hydrogeology, faulting, folding, and fracturing.

Based on the modeled benefits of setbacks for wells near streams and considering existing provisions in the County Code, we are establishing specific setbacks for different tiers of applicants. Tier 1 wells must maintain a minimum 50-foot setback from the stream, and Tier 2 wells are required to maintain a 100-foot setback. Setbacks for Tier 3 wells will be determined based on the criteria necessary to meet stream depletion standards, potentially advancing to Tier 4 standards if compliance is not feasible. Given the modeled potentially significant adverse impacts of stream depletion within 1000 feet (Tier 1) and 2000 feet (Tiers 2 and 3), these wells must adhere to standards aimed at minimizing impacts on streamflow (see Resource Protection Policy), except in cases where a Health Officer designates a stream as exempt.

Detailed Observations Relative to Direct Streamflow Depletion:

- The amount of depletion is moderately reduced by a greater setback from the creek in aquifers characterized by high transmissivity and low storativity. Increasing the setback from 50 ft to 1000 ft reduces the amount of depletion by 25-30% for formations with moderately favorable aquifer properties concerning stream depletion impacts. Conversely, in aquifers with low transmissivity and high storativity, increasing the setback from 50 feet to 1000 feet reduces depletion by approximately 55% for formations with highly favorable aquifer properties.
- 2. Wells pumping 10 af/y or less have very minimal impact on direct flow depletion: less than 0.01-0.02 cfs at a setback of 50 ft from a creek. Incorporating a seal depth of 100 feet further diminishes depletion, with the depletion reduced by

approximately 82% for aquifers characterized by low transmissivity and high storativity values, and depletion reduced by up to 31% for aquifers with high transmissivity and low storativity values. Previous analysis showed that total nonmunicipal pumping has reduced the 10th percentile dry season flow by 2-4% in the Santa Margarita Groundwater Basin and 15-17% in the Mid-County Groundwater Basin. Cumulative impacts are not expected to increase in the future, given the low rate of new rural development and the active management of both basins to reduce the impacts of municipal pumping and raise groundwater levels.

- 3. Pumping from a deeper zone below an aquitard significantly reduces the impact of streamflow depletion, particularly over short time periods (Hunt, 2003). Over modeled 700-day periods, depletion when pumping from below an aquitard with a 50-foot separation is 95-97% less than the depletion at the same distance when pumping from an unconfined aquifer more hydraulically connected to the stream. Over longer periods, such as 10 years, the benefits can remain substantial. With median confining layer hydraulic conductivities, depletion when pumping from below an aquitard with a 100-foot separation is approximately 50% less than the extraction rate. However, under high hydraulic conductivities and low storativity values for both the confining unit and primary aquifer, depletion at a 100-foot separation can still be significant and amount to approximately 80% of the pumped volume. Encouraging new and replacement wells to have a deep seal below an aquitard is expected to be a highly effective strategy for reducing streamflow depletion. These conditions are expected to occur within the Monterey Formation and certain parts of the Purisima Formation.
- 4. Some of the calculations were done assuming the annual volume of pumping all took place in 180 days during the dry season. However, if a 2-year drought was assumed, with the same rate of pumping assumed for the dry season for 700 days, the amount of depletion increased by 17% in the Purisima AA and 56% in the Santa Margarita. If the pumping was from below an aquitard, depletion increased by about 90% in both aquifers when compared with the 180-day scenario, although the amount of depletion was still only 1.6% of the pumping volume.
- 5. Incorporating a deep seal within 1000 feet of a stream is an effective method to mitigate streamflow depletions and reducing drawdown in the upper portion of the aquifer, where the stream is most likely closely interconnected to (Figure 1). This mitigation strategy is particularly impactful for streams connected through aquifers with low permeabilities. However, the degree of reduction in depletion is notably

more pronounced when the well is closer to the stream, likely due to the attenuation of the cone of depression.

For wells with extraction rates of less than 100 AFY located beyond 1000 feet from the stream, the impact of the well seal diminishes (Appendix E) as the curvature of the cone of depression flattens out at farther distances. At these longer distances, the overall drawdown resulting from the pumping volume of the aquifer becomes the primary factor contributing to streamflow depletion.

For instance, when assessing the effects of wells situated 50 feet from a stream, tapping into an aquifer with median values of transmissivity and storativity in the Santa Margarita Formation, a seal depth of 100 feet is projected to decrease stream depletion by approximately 54%, while a 200-foot seal depth could reduce it by around 72%.

For wells positioned 200 feet from the stream under similar geological conditions, a 100-foot seal depth is estimated to mitigate stream depletion by approximately 43%, and a 200-foot seal could reduce it by approximately 62%.

However, when evaluating the impacts of wells situated at farther distances, such as 1000 feet from the stream, the effectiveness of the seal diminishes significantly. In this scenario, with aquifers of similar properties as above, a 100-foot seal depth is anticipated to reduce stream depletion by only 3%, while a 200-foot seal might reduce it by just 5%.



Figure 1- Drawdowns at Different Seal Depths (TTim, Bakker 2013)

6. Beyond 1000 feet, well seal depths are not expected to have a significant impact for wells using less than 100 AFY (see observation #7), and the primary driver to further reduce stream depletion depends on increasing the distance between the stream and the well. For example, considering depletion modeled for wells without seals located in aquifers with high transmissivity and low storativity values, where the zone of influence is expected to be most extensive, stream depletion impacts are reduced by approximately 50% when the well location is increased from 800 feet to 2000 feet (Figure 2). The reduction is projected to be even more significant with distance for aquifers with lower permeabilities.



Figure 2- Stream Depletion Beyond 800 Feet without Seals (TTim, Bakker 2013)

7. Tier 1 wells are expected to have a minimal impact on stream depletion, given their expected requirements, which include a minimum 50-foot stream setback and a 100-foot seal depth when situated in close proximity to the stream. At 50 feet, the maximum estimated depletion ranges from nearly negligible (0.00002 cfs) to 0.0032 cfs (Appendix F with the former corresponding to very low permeable conditions, and the latter corresponding to very permeable conditions. These ranges are projected to be even lower for streams with streambed resistance or scenarios where an aquitard is situated between the stream and the well screen.

Streamflow Depletion Analysis Using USGS Analytical Models:

For the USGS application, three models were primarily used: a partially penetrating stream with nearby pumping from an unconfined aquifer (Hunt, 1999) a partially penetrating stream in an aquitard overlying a pumped aquifer (Hunt, 2003), and a fully penetrating stream with no streambed resistance (Jenkins, 1968). Hunt, 2003 was used to evaluate the effects of requiring a deep seal to the first impermeable layer. Below is a figure showing the set-up for running STRMDEPL08 for pumping from an aquifer associated with the Purisima AA formation with a stream that partially penetrates the aquifer and has streambed resistance (left), and with a stream partially penetrating an impermeable layer with properties similar to the Monterey Formation overlying a pumped aquifer (right). Aquifer parameters are taken from the Groundwater Sustainability Plans, with generally the median figures used (see Table 3).



The USGS analytical models were run for two different aquifer types, the Purisima AA, which has the potential for low to moderate permeability, and the Santa Margarita formation, which has the potential for high permeability. The models were run for various pumping rates and stream setbacks (Table 1). The pumping rates were derived from the annual production (af/y), with a worst-case assumption that the total annual amount is drawn during the typical 6-month dry period (180 days) and maintained at a consistent average amount of continuous pumping to achieve that volume. The volume of pumping for 100 af/y at a 50 ft setback was also considered for situations where pumping occurred below an aquitard, over a 700 day period (2-year drought) and a 10-year period, to understand potential long term effects. However, very long-term effects would normally be mitigated by recharge during normal wet winters.

|--|

Depletion (cfs) with indicated setback from creek (ft) 180 days of pumping, unless noted otherwise

Af/y	summer gpm	pumping cfs	50 ft	100 ft	200 ft	1000 ft
0.5	0.6	0.0014	0.001*	0.001	0.0009	0.0007
2	2.5	0.0056	0.004*	0.004	0.0039	0.003
10	12.6	0.0280	0.0204*	0.0201		0.0149
100	125.7	0.2801	0.2035*			0.1486
100	125.7	0.2801	0.2383*	No aquitaro	d, 700 days	
100	125.7	0.2801	0.2613*	No aquitaro	d, 3650 days	
100	125.7	0.2801	0.0095**	Pumping fr	om below aquita	rd
100	125.7	0.2801	0.0181**	Below aqui	tard, 700 days	
100	125.7	0.2801	0.0388**	Below aqui	tard, 3650 days	
250	314.3	0.7002	0.5765*			0.4288
	1000	2.2282	1.619*		1.547	1.1845

			Depletion (cfs) with indicated setback from creek			
Santa M	largarita (T=3000, S=	1)	(ft), 180 days of pumping, unless noted otherwise			
Af/y	summer gpm	pumping cfs	50 ft	100 ft 200 ft	1000 ft	
0.5	0.6	0.0014	0.0004*			
2	2.5	0.0056	0.0018*	0.0017	0.0012	
10	12.6	0.0280	0.0089*			
20	25.1	0.0560	0.0177*			
50	62.9	0.1400	0.0443*			
100	125.7	0.2801	0.0885*	0.0869 0.0839	0.0616	
100	125.7	0.2801	0.1383*	No aquitard, 700 days		
100	125.7	0.2801	0.1994*	No aquitard, 3650 days		
100	125.7	0.2801	0.0023**	Pumping from below aquitar	d	
100	125.7	0.2801	0.0044**	Below aquitard, 700 days		
100	125.7	0.2801	0.0100**	Below aquitard, 3650 days		
	1000	2.2282	1.1000*	1.0456	0.7798	

*Uses Hunt, 1999 model with a streambed conductance of 1 (ft/day)

**Uses Hunt, 2003 model using aquitard properties similar to the Monterey Formation

Table 1- Key Results Using USGS Models (STRMDEPL08, Reeves 2008)

Analyzing Ranges of Streamflow Depletion and Seal Depth Impacts:

In our analysis of streamflow depletion, we focused on evaluating the upper and lower range of impacts by analyzing various models. Specifically, we examined models that assume a fully penetrating stream without streambed resistance, such as those by Glover, Jenkins, and Bakker (with streamed resistance as an optional parameter). These models predict more significant streamflow depletion compared to models that incorporate streambed resistance or consider partially penetrating streams, such as Hunt's models. Our simulations utilized the aquifer properties of the Santa Margarita Formation under unconfined conditions. This formation was selected because it represents one of the primary water-bearing units in the county, which is also commonly interconnected with surface water. With its potential for high transmissivity/high hydraulic conductivity values and low specific yield values, streams and aquifers associated with the Santa Margarita Formation are particularly susceptible to significant stream depletion (refer to Table 2 for aquifer properties).

We conducted the models for various pumping rates and stream setbacks over a 700-day and 3,650 day periods (Appendix B), corresponding to a 2-year and 10-year drought cycle. During the 2-year timeframe, stream discharge reaches near-equilibrium with unconfined aquifers under steady-state conditions (see Figure 3). To simulate drought conditions and the worst-case effects of intermittent pumping (all water extraction occurring during dry period), we derived pumping rates from annual production, assuming that the total amount is drawn during the typical 6-month dry period and maintained over the drought period. This effectively doubles the amount of typical usage during normal years over the modelled period and serves as a very conservative approach (e.g., 2 AFY wells are modeled as 4 AFY wells).

To analyze the influence of seal depths on stream depletion, we employed the TTim model developed by Bakker in 2013, known for its effectiveness in simulating transient flow in multi-layer systems. The TTim model also served as our primary tool for assessing the worst-case and most extreme impacts on streamflow depletion.

Our simulation environment emulates a homogeneous aquifer divided into three layers, each 100 feet thick. Despite this division, all layers share identical aquifer properties, effectively representing one homogenous aquifer. The top layer is designated as phreatic to mimic unconfined conditions. The simulation includes a well screen positioned sequentially in each layer to assess the impacts of different seal depths for each respective layer. For example, during the third iteration, the well screen is placed in layer 2, effectively simulating sealing of layers 0 and 1. When the iteration has the well screen in Layer 0, the simulation effectively represents no seal for the well. Layer 0 represents the topmost layer (0 - 100 feet below ground surface), while Layer 2 represents the bottommost layer (200 -300 feet below ground surface). The extraction of the well is averaged over the entire screen interval. An example of this simulation is provided in Figure 4, used to assess the worst-

case impacts of a 50 AFY well located 200 feet away from the stream.

Principal Hydrogeologic Unit	Hydraulic Conductivity (feet/day)	Transmissivity (f ee t²/day)	Storativity ¹	Specific Yield ²
Santa Margarita Aquifer Entire Basin	2 – 130	430-7,700	0.008 - 0.02	0.02 - 0.25
Santa Margarita Aquifer Quail Hollow/ Olympia	2 – 50	430 - 6,200	0.008 - 0.02	0.12 - 0.25
Santa Margarita Aquifer Central Portion of Basin	3 – 130	2,000 - 7,700	NA	0.02 - 0.13
Santa Margarita Aquifer Scotts Valley Area	12 – 35	1,000 – 1,700	NA	0.02 - 0.13
Monterey Aquifer ³	0.05 – 6	170 – 1,000	0.00001 - 0.001	0.01 – 0.03
Lompico Aquifer	0.5 – 7	500 - 3,200	0.000001 - 0.001	0.02 - 0.07
Butano Aquifer	0.1 – 6	100 – 1,070	0.000001 - 0.0007	

Table 2-14. Principal Hydrogeologic Units Hydraulic Properties

Adapted from Kennedy/Jenks Consultants (2015); NA = non-applicable given unconfined conditions

¹ Storativity is the volume of water released from confined aquifer storage per unit decline in hydraulic head in the aquifer per unit area of the aquifer.

² Specific yield is the amount of water released from an unconfined aquifer if allowed to drain completely under force of gravity.

³ The Monterey Formation is not a principal aquifer but is included here as there are aquifer test data available for it, and because its occurrence between 2 principal aquifers plays an important role in the hydrogeology of the Basin.

Table 2 "Principal Hydrogeologic Units Hydraulic Properties", (Kennedy/Jenks Consultants, 2015)



Figure 3- Stream Depletion Over 50 Years (streamDepletr, Li et. al)



Figure 4-Simulation of Well Seal Depth Impacts on Groundwater Extraction at Different Depths (TTim, Bakker 2013)

Tool Selection for Applicants:

In evaluating streamflow depletion due to groundwater pumping, county staff have used numeric groundwater models where they have been developed for the Mid County and Santa Margarita groundwater basins. Staff have also applied the analytical models developed by Hunt, Jenkins, Li, and Bakker. Staff have assessed more complex models cited by Li, et.al. and Bakker, recognizing their significance and usefulness in establishing thresholds for policy development and testing. These models are particularly valuable for evaluating impacts over extended timeframes, intermittent pumping, seal depths, setbacks, and areas requiring more thorough analysis.

While the County staff found these programming models (Li. et al, Bakker)useful, they did not observe significant differences in the fundamental calculation for stream depletion (without incorporating well seals) when assuming fully penetrating streams with no streambed resistance compared to the simpler USGS web-based application, especially when analyzing single-point scenarios that focus on streams closest to the well. While the USGS web-based application may be suitable for Tier 3 applications, Tier 4 applicants must prepare a report by a professional geologist, engineering geologist, or professional engineer to evaluate more detailed projected impacts, including the cumulative effects on streamflow in the overall basin. Because of this requirement, we encourage these consultants to consider using more advanced tools, particularly Li et al. for evaluating cumulative impacts on a network of streams and Bakker for evaluating the influence of deeper seals in minimizing stream depletion impacts.

Local Aquifer Properties: Range (typical value used)	Transmissivity (ft^2/day) {gpd/ft}	Storage/ Storativity	Specific Yield	Hydraulic Conductivity
TP-a - Purissima A	(2000) {15,000}	0.00055	0.02-0.07 (0.05)	5.2
TP-aa- Purissima AA	(600) {4500}	0.03100	(0.02)	1.7
TSM - Santa Margarita	430-7700 (3000) {22,500}	0.01	0.02-0.25 (0.2)	2-130
TLO - Lompico	500-3200 (2000) {15,000}	0.0000020	.0207 (.05)	0.5-7
Aromas/Purisima F	(4000) {30,000}	0.004		
Tm-Monterey	170-1000	0.00001-0.001	.0103	.056

Table 3- Aquifer parameters from Groundwater Sustainability Plans

		Dry Season I	Flows, cfs	(All Years)	
		10th		90th	
Creek		Percentile	Median	Percentile	Source
	Estimated Natural Flow*	0.509	1.08	1.89	FF model*
Poon Cr. @ Mt	Observed *	1.9	2.25	2.82	FF Database*
Hormon Pd	Est.depletion by total gw pumping	0.5	0.5	0.5	GSP model
	% depletion**	21%	18%	15%	
(0303)	Est depletion by Non-Mun pumping	0.08	0.08	0.08	Apply Basin-wide proportion from GSP Model
	% Non-muni depletion	4%	3%	3%	
	Estimated Natural Flow*	15.2	20.2	23.7	FF model*
Contoronzo	Observed*	12	19	32	FF Database*
San Lorenzo River @ Big Trees (USGS)	Est.depletion by total gw pumping	1.5	1.5	1.5	GSP model
	% depletion**	10%	7%	4%	
	Est depletion by Non-Mun gw pumping	0.23	0.23	0.23	Apply Basin-wide proportion from GSP Model
	% Non-muni depletion	2%	1%	1%	
	Estimated Natural Flow*	0.0542	0.153	0.452	FF model*
Maara Cr	Observed	0.15	0.3	0.5	Estimated based on Occasional Measurements
MODIECI	Est.depletion by Non-Mun gw pumping	0.03	0.03	0.03	Water Budget
	% depletion	17%	9%	6%	
Soquel Cr. @ Soquel (USGS) ***	Estimated Natural Flow*	2.44	3.05	5.28	FF model*
	Observed *	0.84	2.86	8.05	FF Database*
	Est.depletion by total gw pumping***	1.4	1.4	1.4	GSP model
	% depletion	57%	33%	15%	
	Est depletion by Non-Mun pumping	0.15	0.15	0.15	GSP Model
	% Non-muni depletion	15%	5%	2%	

Estimated Surface Water Depletion from Groundwater Pumping in Selected Santa Cruz County Streams

Notes

* Estimated Natural Flow and Observed Flow is provided by the California Unimpaired Flow Database, v2.1.2 (Zimmerman, et.al., 2023)

** % depletion is the estimated depletion divided by the greater of the estimated natural flow, or the observed flow plus the estimated depletion *** Soquel Creek experiences significant riparian surface diversions, potentially 0.5-0.7 cfs (RCDSCC, 2019).

The potential effect of surface diversions has not been factored into this table, other than where the estimated natural flow is used.

Table 4- Estimated Natural Flows and Depletion Based on Natural Flows Database, Streamflow Measurements, Local Groundwater Modelling, and Water Budgets

Well Interference

Staff have used the Modified Theis Non-Equilibrium Equation to estimate the amount of drawdown at various distances from a proposed pumping well in order to evaluate the potential for well interference and potential impacts on nearby wells. Values for local aquifer properties, pumping rates and potential setbacks were entered in the formula to produce an estimated drawdown. The following table shows the setbacks required for

particular pumping rates in order to keep the drawdown less than 5 ft after 180 days of pumping.

Pumping Rate (GPM)	2	8	20	50	100
Aquifer					
TP-a/TLO	10	10	10	10	150
TP-aa	10	10	25	500	1400
TSM	10	10	10	10	25

			Input	
Equation	s=(264Q/T)*log(.3Tt/((r^2)S)	Values	Result
Q	Discharge	gpm	50	
Т	Transmissivity	gpd/ft;(7.48*ft^2/d)	4500	
S	Storage Coefficient	dimensionless	0.020	
t	Pumping time	days	180	
r	Distance	ft	100	
s=	drawdown-calculated	ft		9.0

Staff is proposing to use a standard of 50 ft separation for de minimis wells and replacement non-de minimis wells, although a greater setback could be required for new non-de minimis wells after applying the Modified Theis Non Equilibrium Equation to the specific well and aquifer properties.













Stream depletion (cfs) after 10 years: -0.0035



Stream depletion (cfs) after 10 years: -0.0100





Stream depletion (cfs) after 10 years: -0.0439





Stream depletion (cfs) after 10 years: -0.0721



Stream depletion (cfs) after 10 years: -0.0879



Stream depletion (cfs) after 50 years: -0.0427



Stream depletion (cfs) after 50 years: -0.0832



Stream depletion (cfs) after 50 years: -0.0945











Stream depletion (cfs) after 700 days: -0.1572


Stream depletion (cfs) after 700 days: -0.1505



Stream depletion (cfs) after 700 days: -0.1422



Stream depletion (cfs) after 700 days: -0.1334



Stream depletion (cfs) after 700 days: -0.1248



Stream depletion (cfs) after 700 days: -0.1167



Stream depletion (cfs) after 700 days: -0.1092









Stream depletion (cfs) after 700 days: -0.0002



Stream depletion (cfs) after 700 days: -0.0023



Stream depletion (cfs) after 700 days: -0.0032







Stream depletion (cfs) after 10 years: -0.0035



Stream depletion (cfs) after 700 days: -0.0002





Stream depletion (cfs) after 700 days: -0.0023





Stream depletion (cfs) after 700 days: -0.0032





Stream depletion (cfs) after 700 days: -0.0058









Stream depletion (cfs) after 700 days: -0.0796





Stream depletion (cfs) after 700 days: -0.0115





Stream depletion (cfs) after 700 days: -0.1125





Stream depletion (cfs) after 700 days: -0.1594





Stream depletion (cfs) after 700 days: -0.0001





Stream depletion (cfs) after 700 days: -0.0020




Stream depletion (cfs) after 700 days: -0.0030

















Stream depletion (cfs) after 700 days: -0.0074





Stream depletion (cfs) after 700 days: -0.1009





Stream depletion (cfs) after 700 days: -0.1505





Stream depletion (cfs) after 700 days: -0.0000





Stream depletion (cfs) after 700 days: -0.0006





Stream depletion (cfs) after 700 days: -0.0018









Stream depletion (cfs) after 700 days: -0.0153









Stream depletion (cfs) after 700 days: -0.0004





Stream depletion (cfs) after 700 days: -0.0306





